

Evaluation of modeled diurnally varying sea surface temperatures and the influence of surface winds

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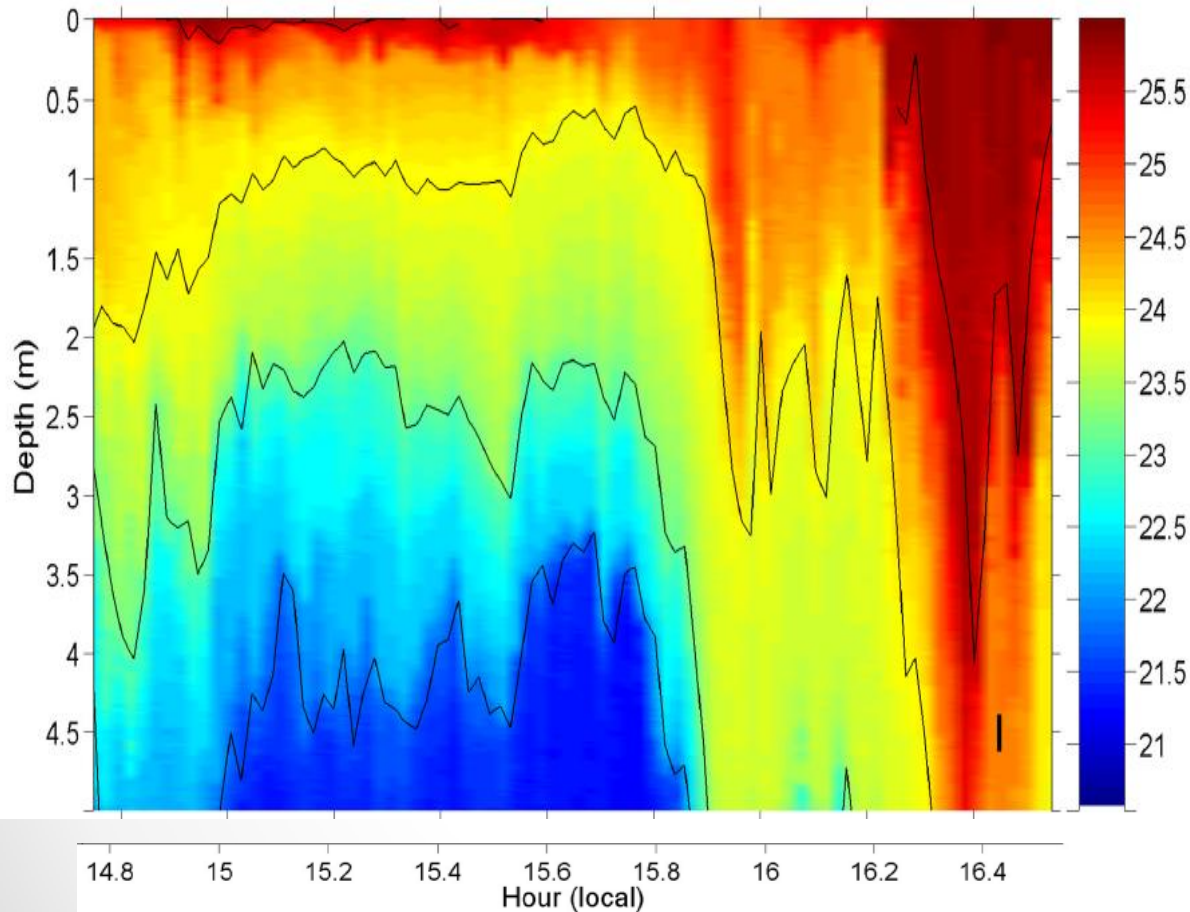


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SSTs exhibit diurnal variability



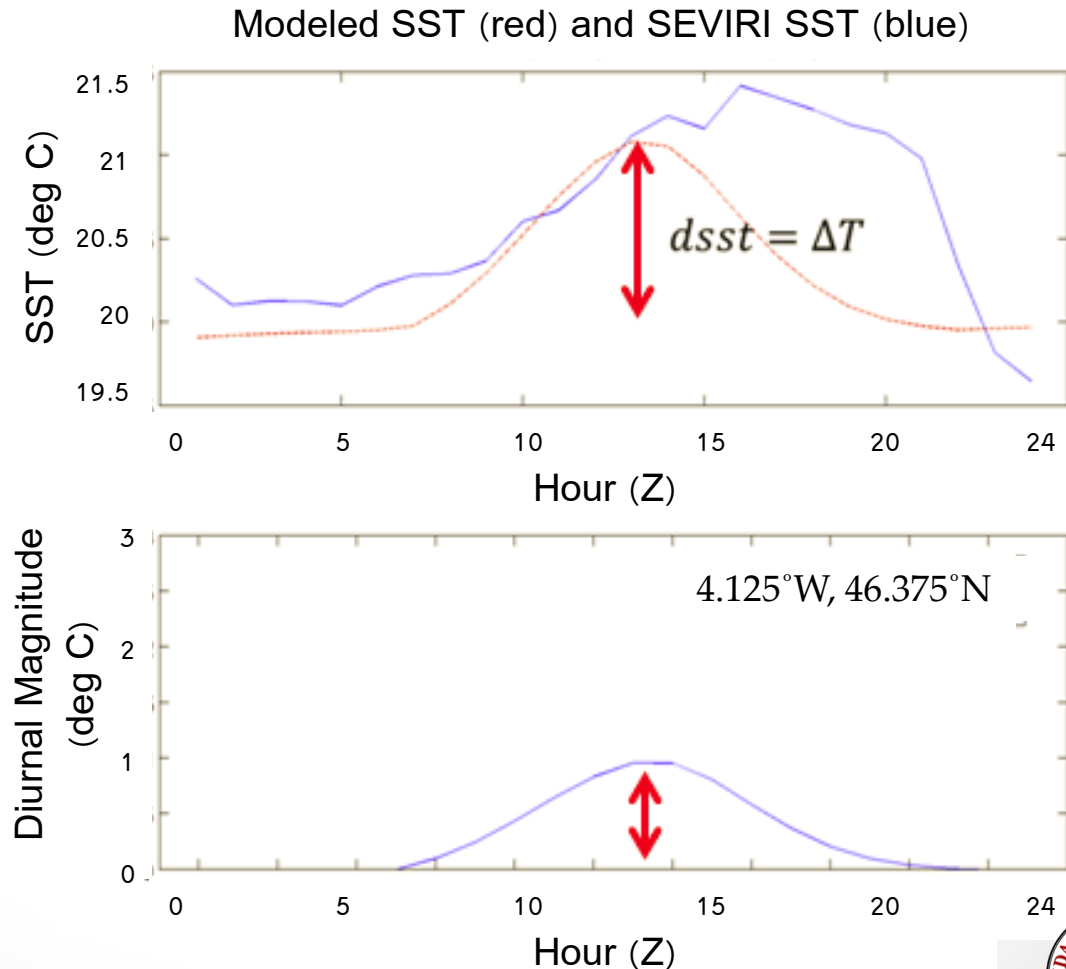
From Ward 2006

- As far as last century, studies have noted ΔT between different depths in upper few meters over course of day
- ΔT can be on the order of 3°C
- Presence of diurnal warming in certain areas can last throughout the majority of daylight hours.



Development of warm surface layer

- Winds are relatively calm (low mechanical turbulence and mixing)
- Ample sunlight absorption into the ocean
- Diurnal cycle observed using model and satellite data



Overview of Research

- Creation of diurnally varying SST dataset using a physical-empirical hybrid model (2000-2004)
- Characterization of diurnal warming in regions of the ocean
 - Diurnal streaks
 - Temporal persistence in Tropics vs. Midlatitudes
- Examine the influence on surface latent and sensible heat fluxes over seasonal time scales
- Analyze sensitivities in model with regard to winds, solar radiation at the surface, and precipitation
- Model evaluation using remotely sensed skin temperature



dSST Product

- **POSH – Profiles in Oceanic Surface Heating Model**
 - Gentemann et al. (2009)
 - Collection of improvements of TOGA-COARE bulk flux algorithm v2.5
 - Fairall et al. (1996)
- **Calculates diurnal warming in a 1-D model at each hour on 0.25° x 0.25° grid**
 - Function of bulk SST and accumulated heat/momentum at each time step.
 - Advection is ignored
 - Ice is excluded
- **Variety of diurnal warming models**
 - Physically based model that captures diurnal warming will likely increase understanding of physical processes of the upper ocean
- **Produced 5 year long dataset (2000-2004)**
 - SST and dSST (magnitude)
 - sensible and latent heat flux with and without a diurnally varying SST
 - using BVW model (Bourassa 2006)



Input Data

- **POSH needs global continuous time series of hourly data**
 - atmospheric bulk variables, radiation, and SST_{bulk}
 - U_{10} , V_{10} , Pressure, T_{10} , Q_{10} , SWR, LWR, Precip Rate
- **NASA's MERRA Reanalysis**
 - Provides hourly data on $1/2^\circ \times 2/3^\circ$ grid for entire globe – regridded to $1/4^\circ$
- **Bulk SST**
 - Need SST closer to SST_{fnd} (Donlon et. al 2007) unaffected by diurnal heating for total effect
 - Reynolds OI daily SST AVHRR-only (Reynolds et al. 2003)
 - $1/4^\circ \times 1/4^\circ$
 - Uses satellite and in situ measurements for blended product
 - Bias adjusted over seven-day period – eliminates diurnal variability
 - Set to 19 m - theoretical maximum dSST thickness (Gentemann personal communication 2011)



Scaled dSST

- Diurnal warming magnitude
- Diurnal warm layer thickness:

$$dSST(z) = e^{-9.5\left(\frac{z}{D_T}\right)^a}$$

$$D_T = \sqrt{\frac{2R_i c_p}{\alpha g \rho}} \frac{\tau_{ac}}{\sqrt{Q_{ac}}}$$

α water thermal expansion coefficient
 ρ density of water
 R_i Bulk Richardson Number
 c_p Heat Capacity of Water
 g Gravitational Acceleration

Wind Speed	a
≤1.5	2
3	3
4.5	5
6	7
≥7.5	9

Requires bulk Richardson number = 0.65
 Assuming a linear temperature profile
 Function of accumulated wind stress and accumulated heat

$$\tau_{ac}(t) = (1 - \varepsilon_\tau \Delta t) \tau_{ac}(t - \Delta t) + \tau_{tot} \Delta t$$

$$Q_{ac}(t) = (1 - \varepsilon_\chi \Delta t) Q_{ac}(t - \Delta t) + Q_{tot} \Delta t$$

Dissipation coefficients make diurnal warm layer “interact” with mixed layer and lose heat from bottom DWL

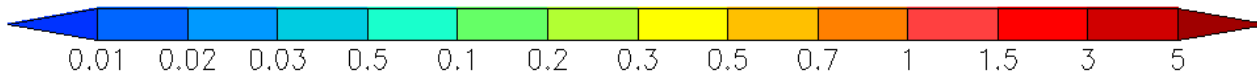
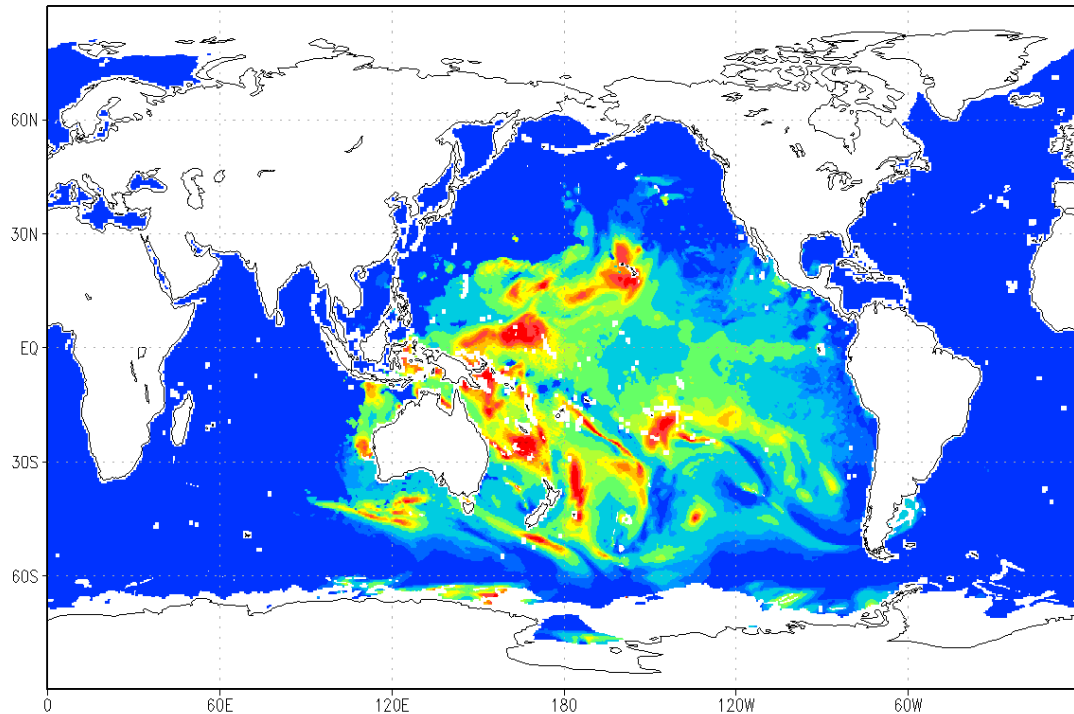
Salinity dependence is not explicit in model

From Gentemann et al. 2009



dSST Subdiurnal Spatial Structure

dSST mag (deg C) 2000 01 01 1:00 Z



Diurnal warming magnitude ($^{\circ}\text{C}$)

- Sun synchronous
- Occur at almost all latitudes
- Reflect current meteorological state
 - Cyclones
 - Fair weather features
- Non-linear non-negligible increase in surface heat fluxes

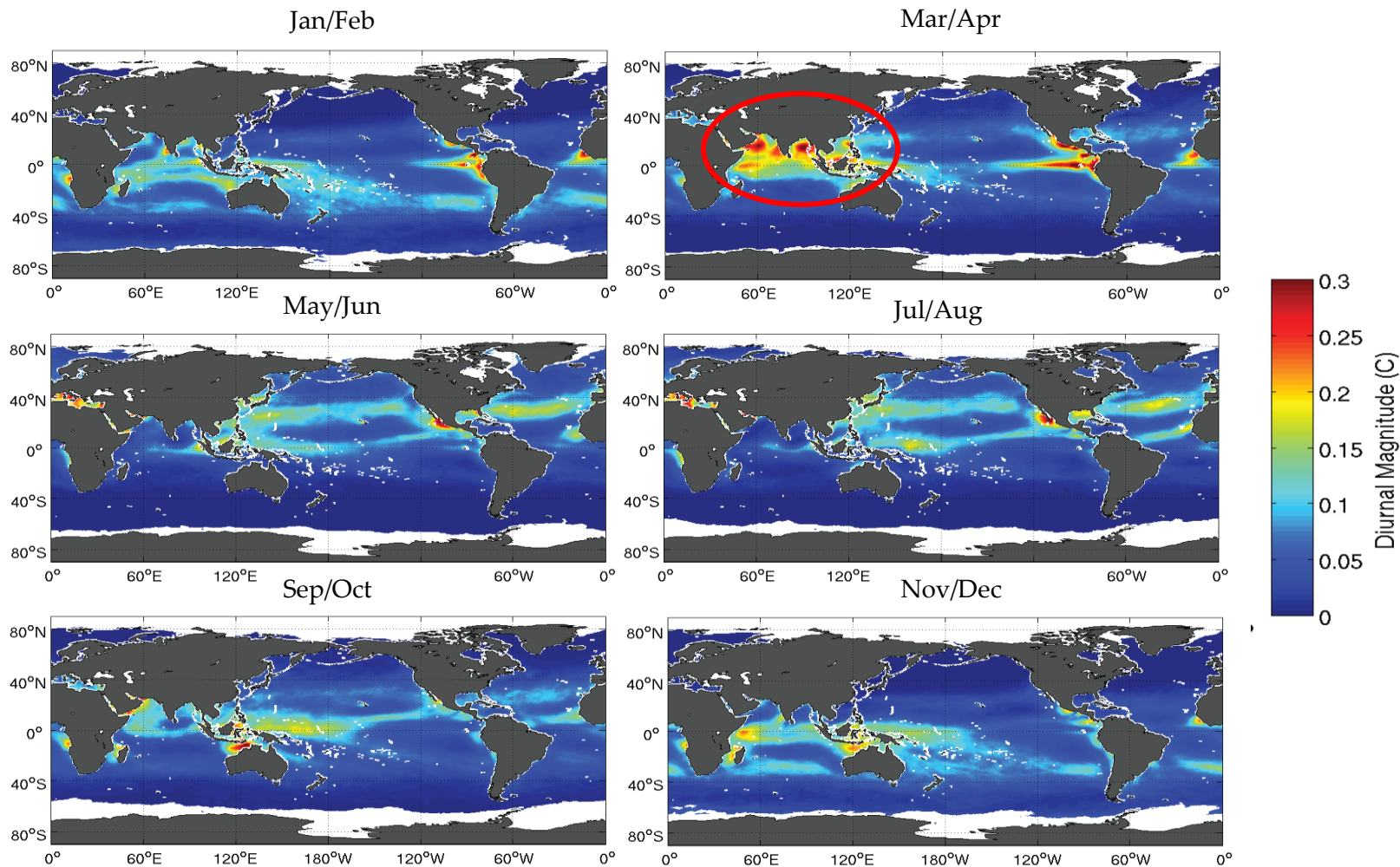


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Bimonthly Average dSST



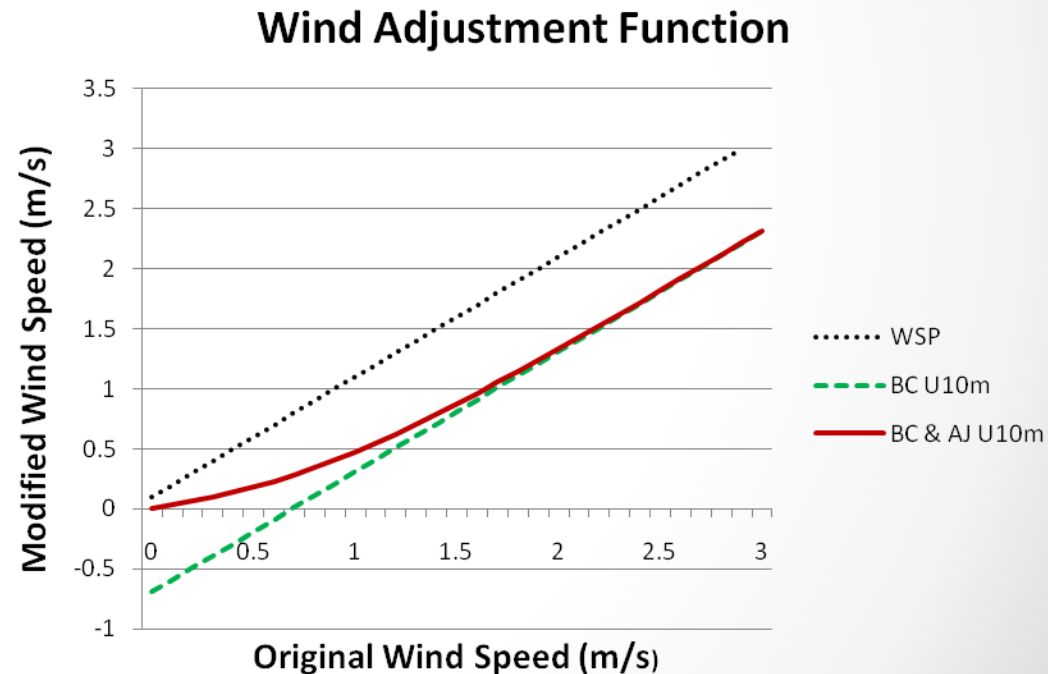
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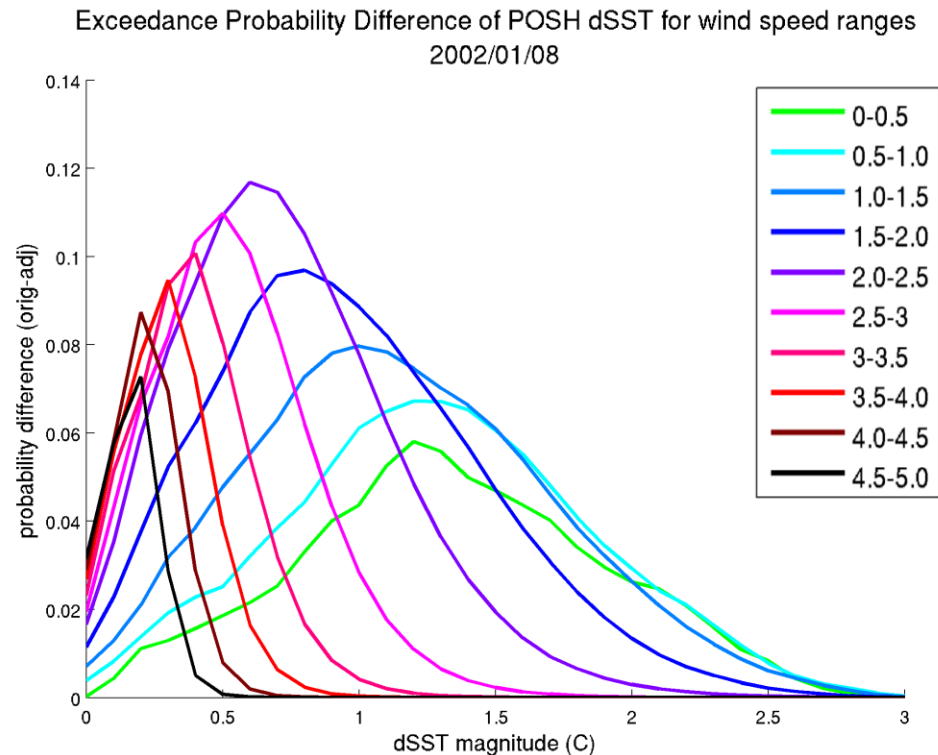
Effect of Potential Biases in Winds

- We can estimate the error associated with biases in the wind speed data
- Roberts et al. [2010] compared MERRA with in situ measurements
 - MERRA had positive bias for entire distribution of winds
- Best guess estimate of systematic bias in winds is 0.69 m/s



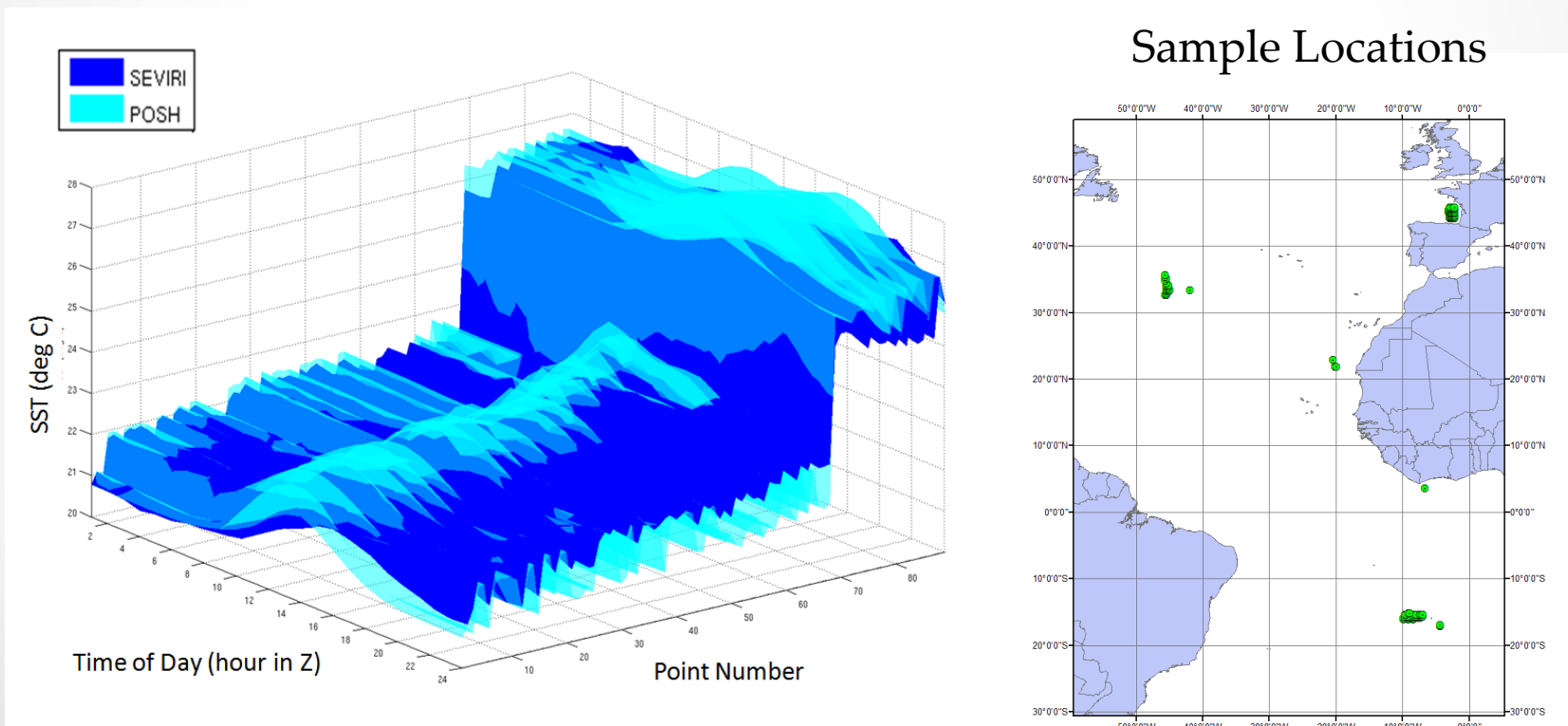
Effect of Potential Biases in Winds cont.

- Potential biases affect the short-term diurnal warming – exceedance curves
- Differences are +
 - Increased likelihood of dSSTs at all wind speed ranges
- Largest probability increase between winds of 2.0 – 2.5 m/s
 - Lowest wind speed with the largest bias correction value
- Isolated events can be as large as 1°C
- Bimonthly differences are less than 0.06°C (not shown)

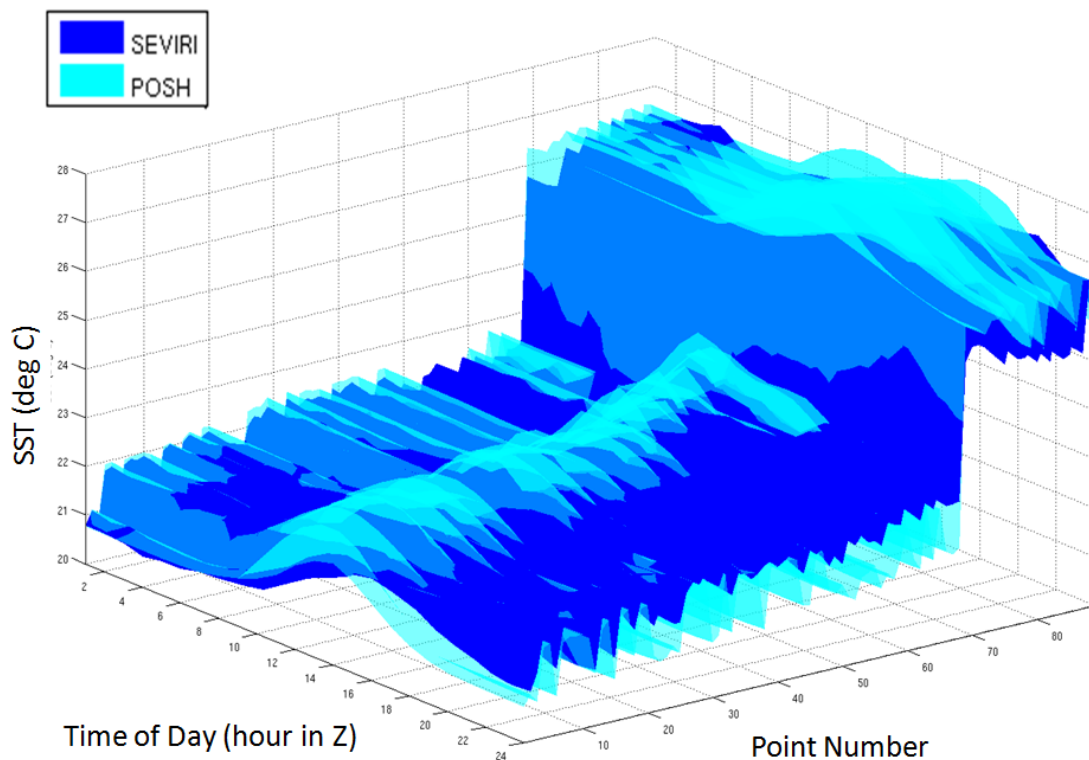


Time Series of Diurnal Cycles

- Using the best sampled data points on a given day, compare modeled SST to SEVIRI SST at each hour



Time Series of Diurnal Cycles



- Reynolds SST estimates foundation temperature reasonably well
- POSH consistently overestimates peak but dissipates heat too quickly for smaller peak SST
 - Tuning of dissipation coefficients for different regions of ocean may prove useful

Conclusions

- The diurnal cycle of sea surface temperatures is modeled using Gentemann et al.'s POSH model (2009)– a physical-empirical hybrid model based on TOGA-COARE flux model
 - exponential adjustment to the diurnal warm layer thickness;
 - a function of accumulated wind stress and heat adjusted by dissipation coefficients to interact with mixed layer
- Occurs throughout the world's ocean
 - except in areas of ice,
 - sometimes in thousand kilometer streaks
 - Non-linearly affects surface fluxes
- Best guest estimate of bias correction (0.69 m/s) influenced dSST on short timescales but less so on semi-seasonal scales
 - Sensitive to bias adjustment at low wind speeds
- Using regrided SEVIRI SST swath data, selected points based on their sampling were selected to use in a comparison under a best case scenario approach.
- The Reynolds SST was representative of a foundation temperature while the peaks in POSH were consistently overestimated



Future Work

- Couple diurnally varying SSTs to boundary layer model
 - Input from satellite winds, SSTs
 - Varying surface fluxes
 - Potentially with HYCOM
- Examine effects
 - Divergence, wind feedbacks on SST



References

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- [7] Bourassa, M. A., 2006: Satellite-based observations of surface turbulent stress during severe weather. *Atmosphere-Ocean Interactions*, Volume 2, W. Perrie, Ed., *Advances in Fluid Mechanics*, Vol. 39, WIT Press, 35-52.
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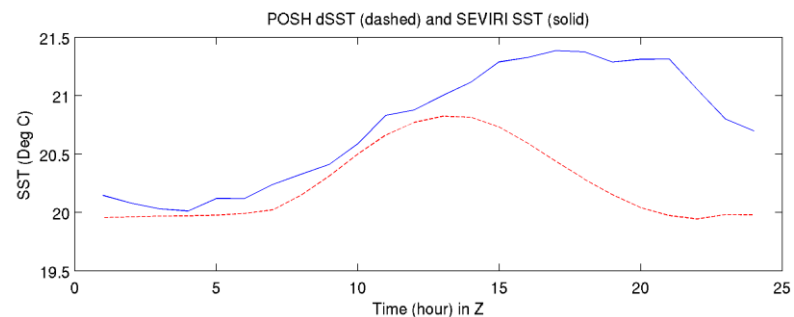


Questions?

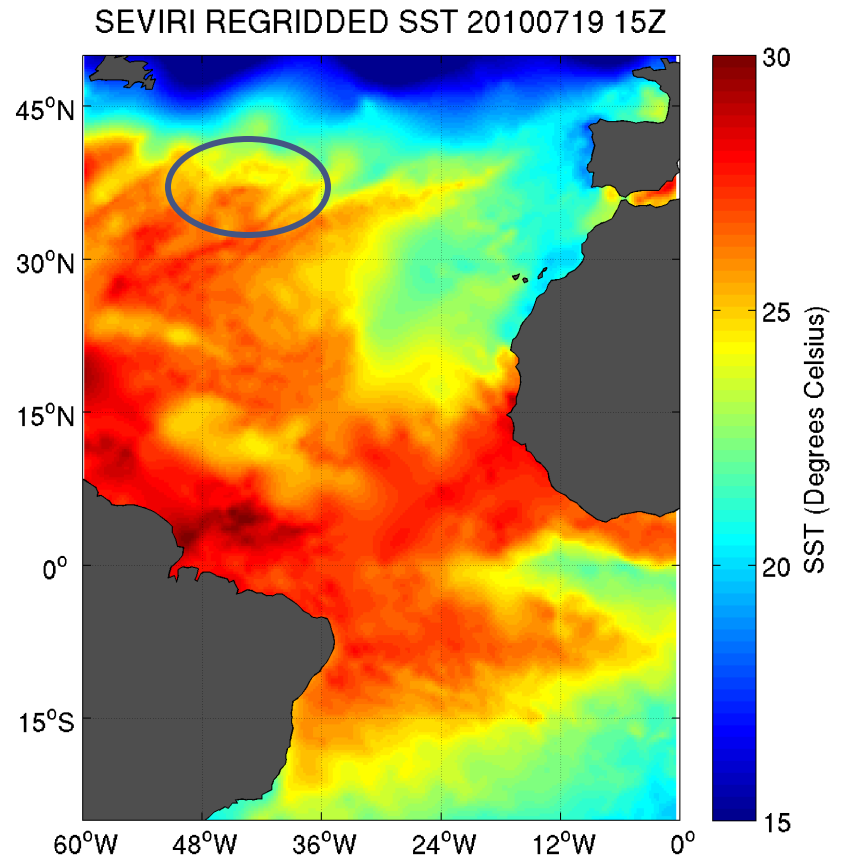
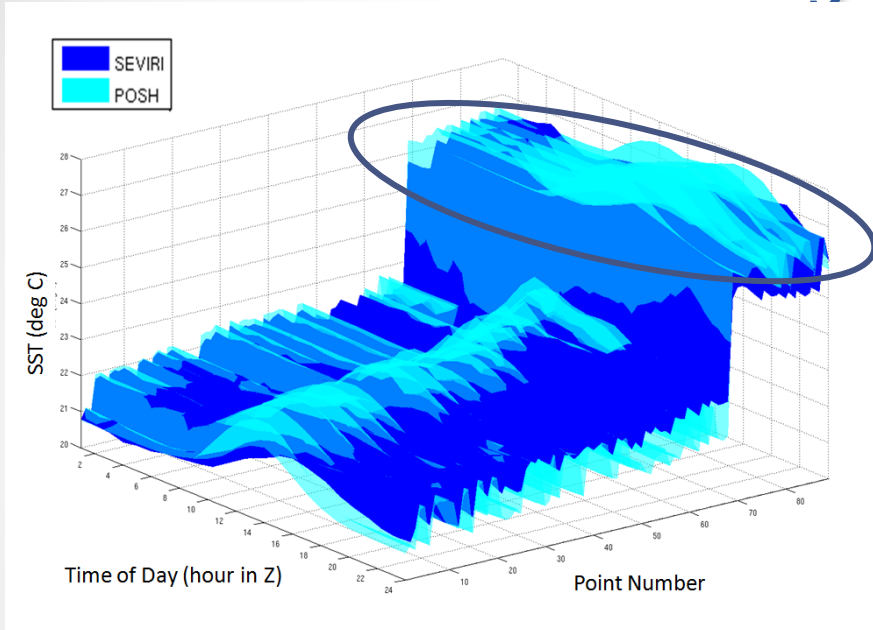
Back up slides

Comparison of Modeled Diurnal Cycle to Measurements

- Want an initial assessment of how well POSH represents the diurnal cycle of global SSTs
- Perform a case study
 - Using IFREMER's MSG/SEVIRI hourly skin SST swath data
 - One of the newer instruments; has already been shown to observe diurnal cycle (Le Borgne et al., 2012)
 - Average underestimation of peak dSST by 0.12 K
- Rerun POSH model to match data availability
 - Sample day in July, in Atlantic Ocean
- Differences are function of
 - Foundation offset
 - Shift in onset of warming
 - Hourly/peak offset



Time Series of Diurnal Cycles



- Areas where SST gradients are large are ambiguous
 - Advection (or large turbulent mixing) may inhibit formation of a classical diurnal warm layer



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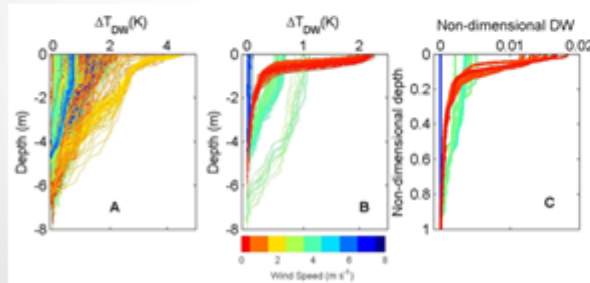
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From [Gentemann et al. 2009](#)

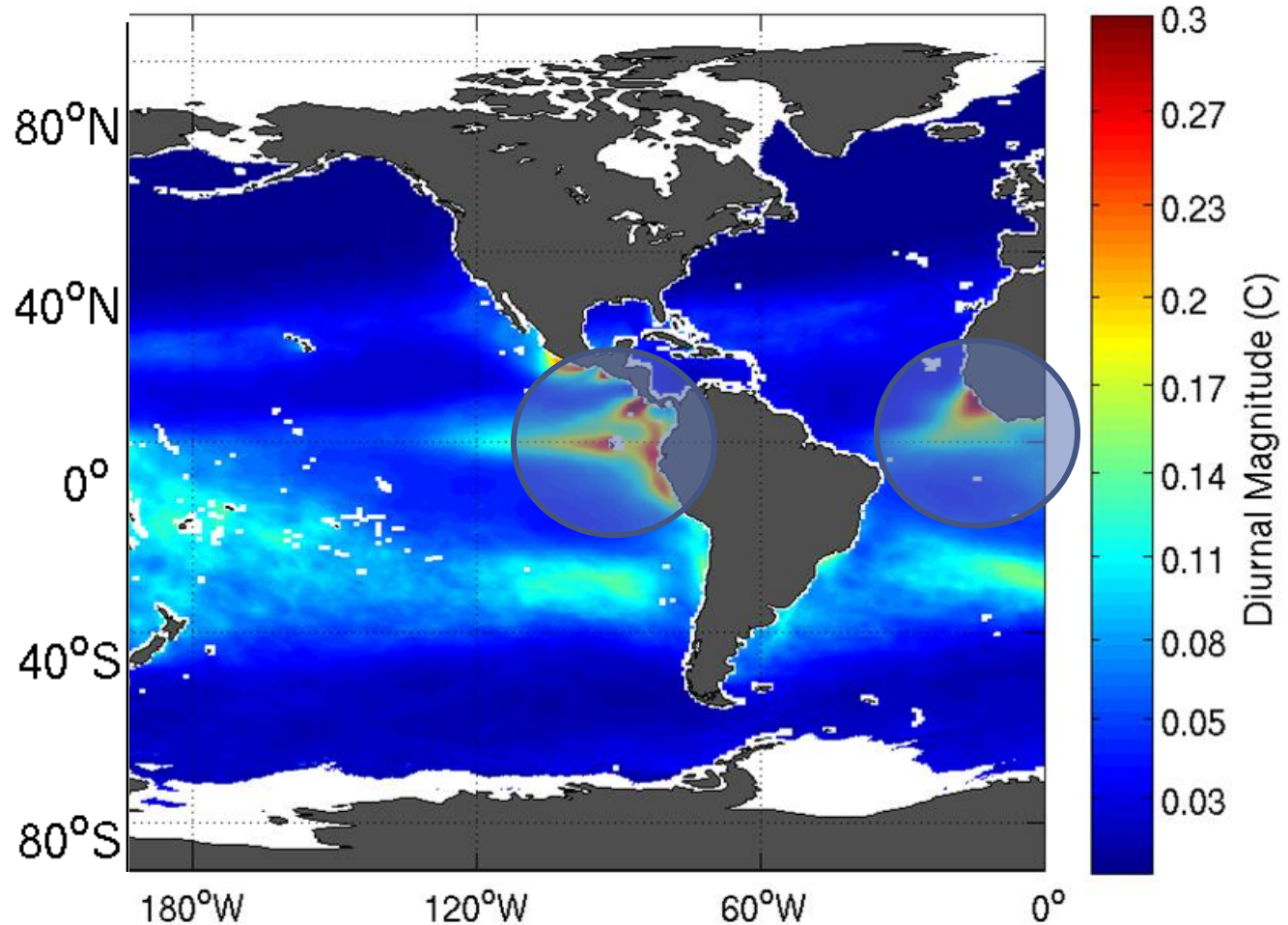
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Bimonthly average dSST

Jan/Feb Composite

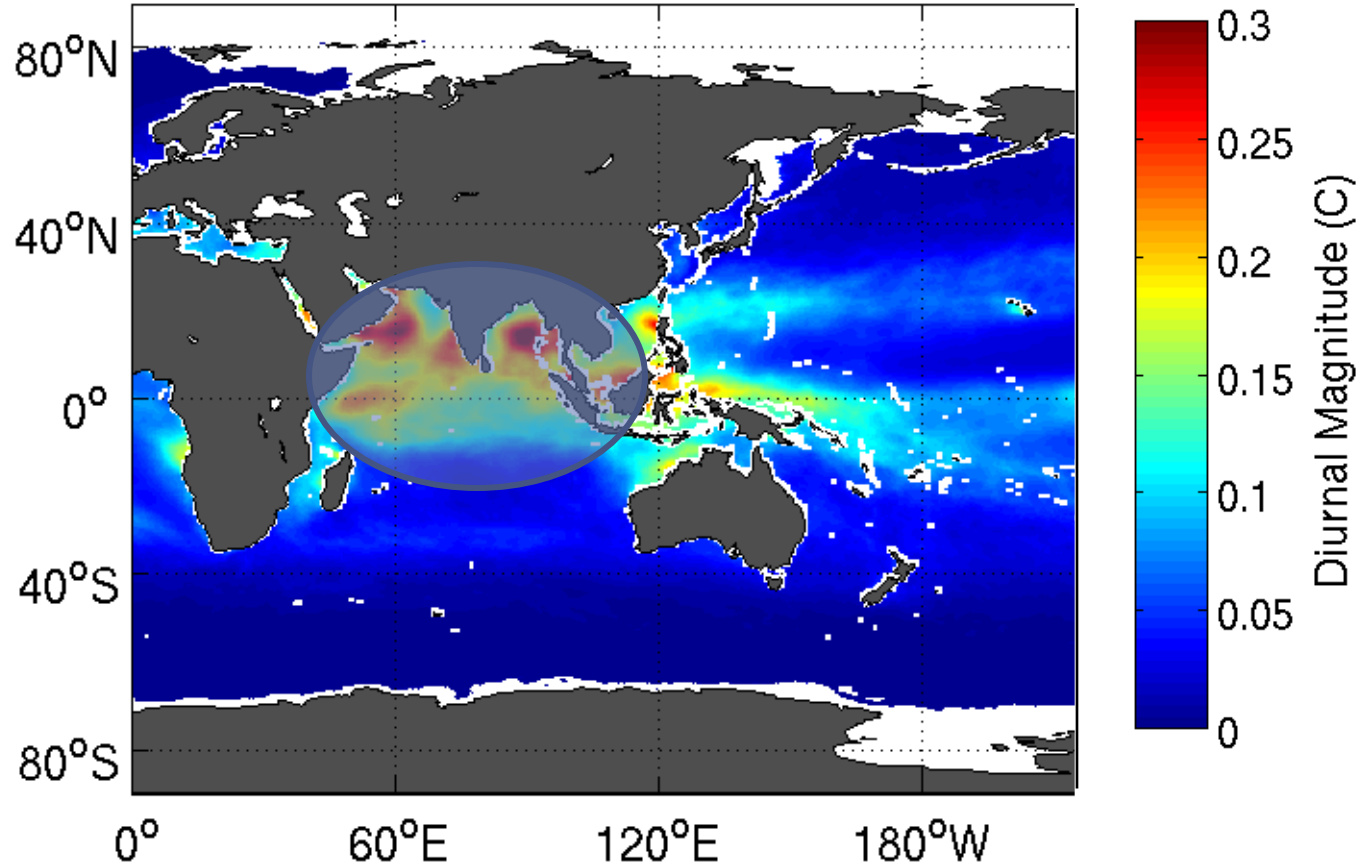
- Low level convergence near the ITCZ



Bimonthly average dSST

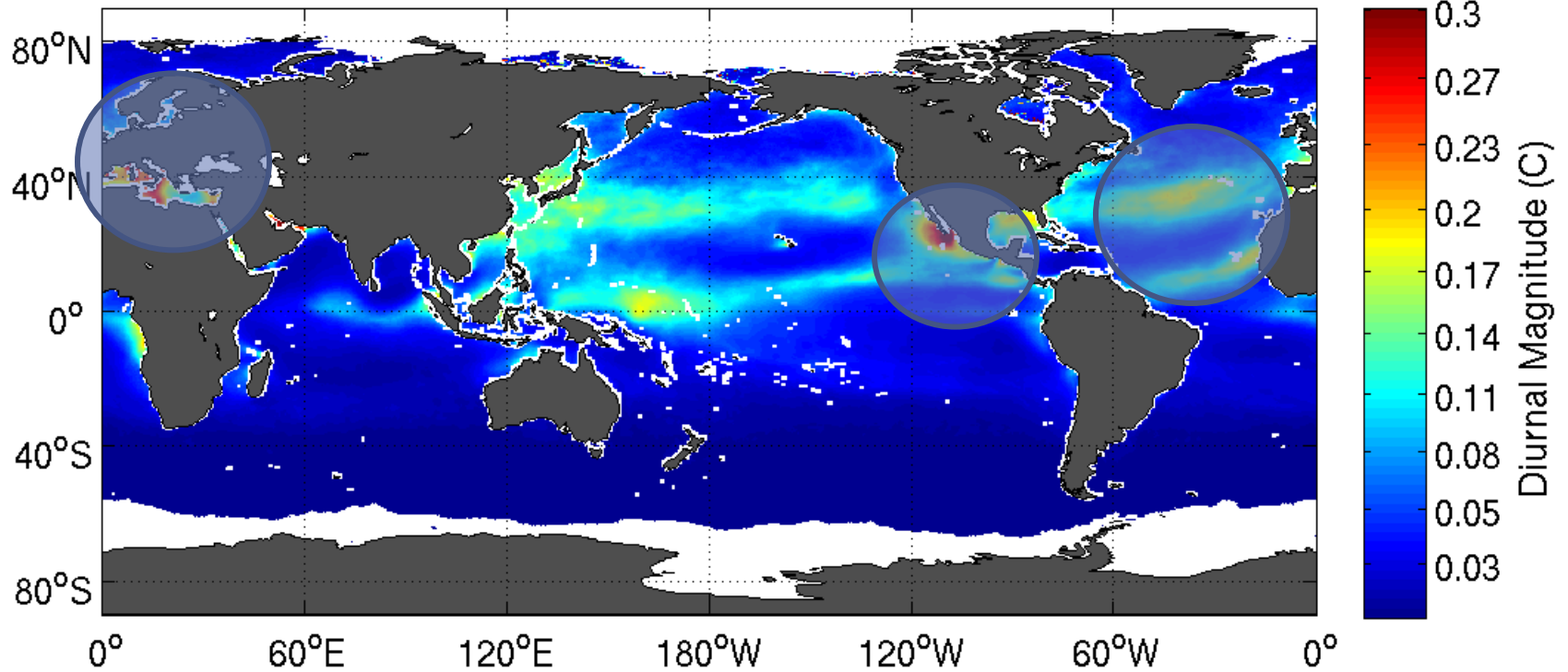
Mar/Apr Composite

- Transitional phase of monsoonal winds



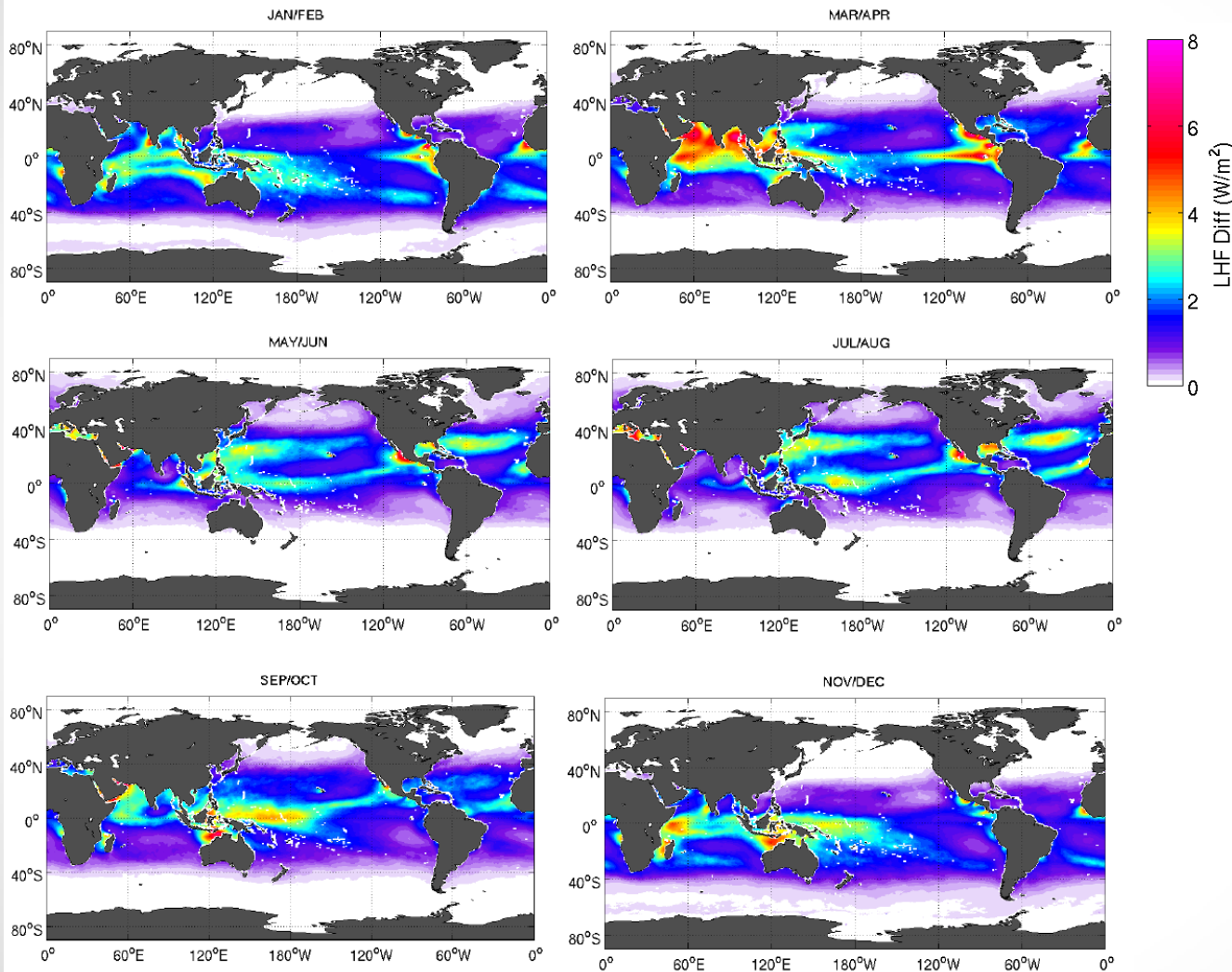
Bimonthly average dSST

Jul/Aug Composite



- Mountain Barriers
- Mediterranean Sea
- Azores High

Bimonthly average LHF difference

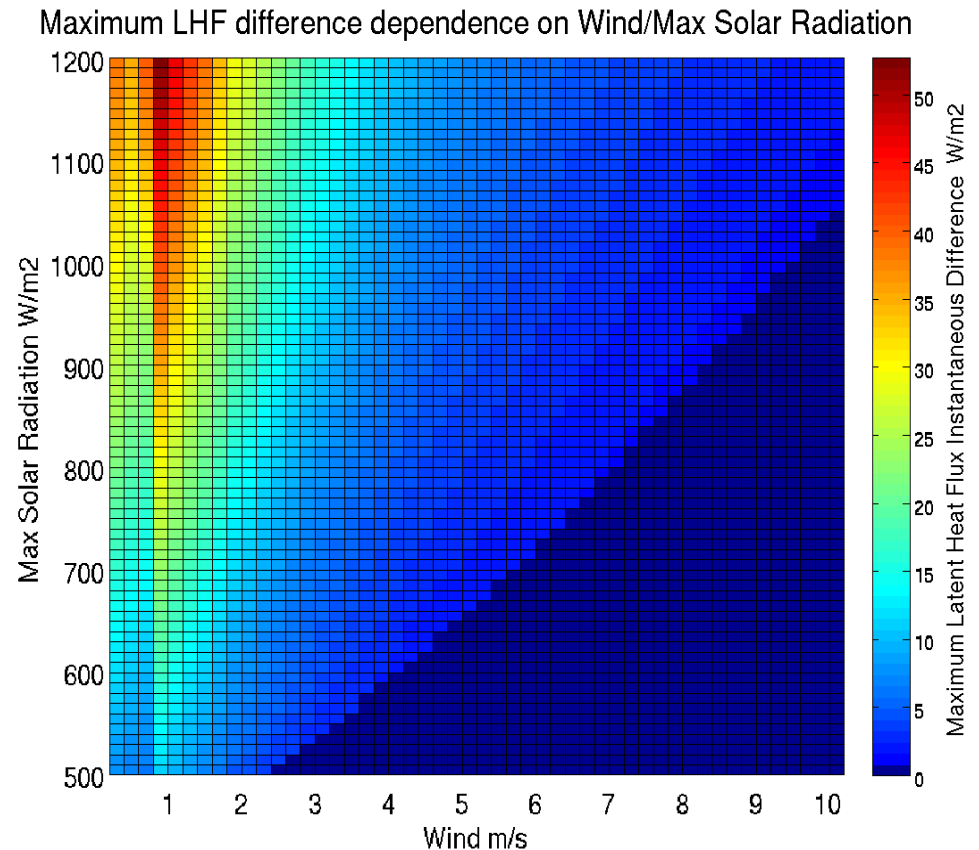


- Latent heat flux accounts for **2/3 of the total net flux^{*}**.
- Heating released to the atmosphere of this magnitude is **considered non-negligible** in terms of response^{*}.
- Supports theory that **diurnal warming can affect convection** in the tropics.

[*Kawai and Wada, 2007; [^]Bellenger and Duvel, 2009]

Flux Pt Sensitivity to Wind/SWR ranges

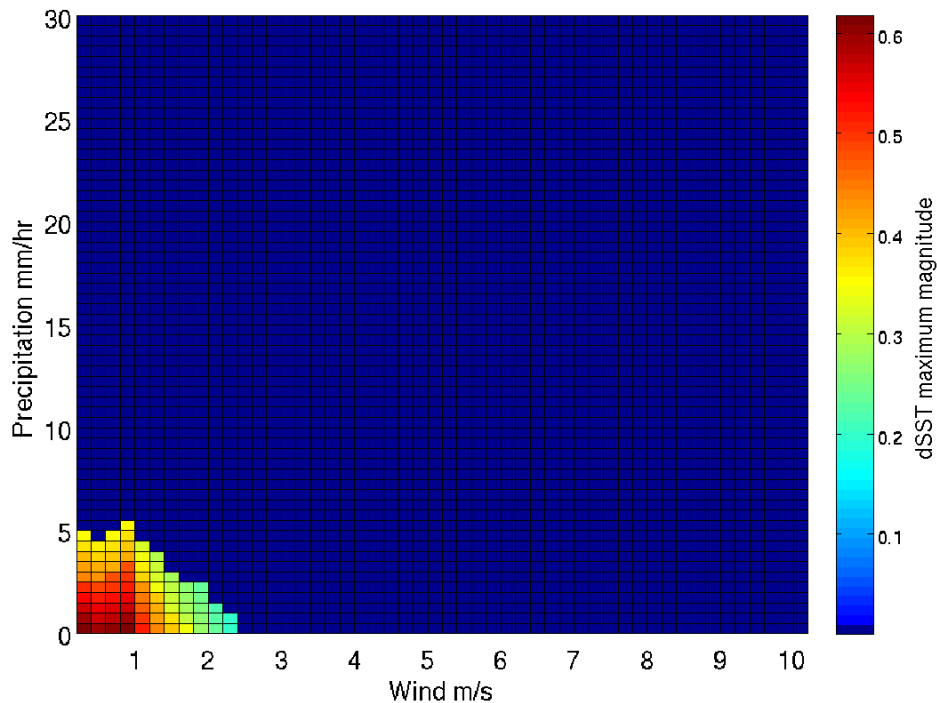
- Flux value at dSST maximum
 - LHF differences expected to increase at dSST increases
 - At wind speed < 0.8 m/s, LHF difference change slows as winds decrease
 - Balance between humidity and wind.
 - Humidity change is less influential to LHF than wind speed.
- Possible consequence largest dSSTs may not correspond to largest LHF deficit.



Max dSST sensitivity to Wind/Precip Ranges

- Similar sensitivity study as Wind/Solar Radiation
 - Precipitation total 0.0 – 30 mm/hr
 - SWR peak is set at 500 W/m² and 700 W/m² for cloudy, rainy conditions

Maximum dsst dependence on Wind/Precipitation
Max Radiation 500 W/m²



Maximum dsst dependence on Wind/Precipitation
Max Radiation 700 W/m²

